

Reply to the discussion by Zhang and Lytton on “A new modelling approach for unsaturated soils using independent stress variables”¹

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We would like to thank the discussers for their interest in our paper and also for their challenging questions, which show their thorough knowledge on the subject. Their comments have enhanced our understanding of the subject area and hopefully may lead to further constitutive model developments.

Question 1: stress-path dependency within the elastic zone

The stress-path dependent elastic response will lead to

$$[1] \quad \Delta(\ln v)|_{ABD} = -\kappa_{vp} \ln \frac{p+s}{p_0+s} + \frac{\kappa_{vp}(s_{sa}+1)}{p_0-1} \left(\ln \frac{p_0+s}{s+1} - \ln \frac{p_0+s_0}{s_0+1} \right)$$

The volume change for stress path ACD is correctly derived as

$$[2] \quad \Delta(\ln v)|_{ACD} = -\kappa_{vp} \ln \frac{p+s_0}{p_0+s_0} + \frac{\kappa_{vp}(s_{sa}+1)}{p-1} \left(\ln \frac{p+s}{s+1} - \ln \frac{p+s_0}{s_0+1} \right)$$

Equation [1] is indeed different from eq. [2], but the actual difference is very small. A numerical example is given as follows. Some qualitative remarks are warranted. When suction is less than the saturation suction, the SFG model is stress-path independent; when suction is greater than but close to the saturation suction, κ_{vs} is close to κ_{vp} , and hence the volume-change difference between stress paths ABD and ACD is small. On the other hand, when suction is very large, the volume change ($-\kappa_{vp} \frac{dp}{p+s}$) caused by a stress

hysteretic behaviour. As such, a loading path ABD followed by an unloading path DCA in Fig. 1 of the discussion would not lead to zero volumetric strain. This hysteretic behaviour constitutes a restriction in classical elastoplasticity theory. We have carefully considered all possible options in the literature and find that the SFG model still provides a better solution than most other options.

It should be noted that the equation derived in the discussion for stress path ABD has a small misprint, and the correct equation should read as

change is close to zero (because of the term $(p+s)$), and the volume change caused by a suction change is also close to zero (because of κ_{vs}). Therefore, the difference in the volume changes between the two stress paths is also small.

Option 1: other models in terms of net stress

Let us now give consideration to other alternatives. The following elastic model is commonly used in the literature (also used in Zhang and Lytton 2007):

$$[3] \quad dv^e = -\kappa_{vp} \frac{dp}{p} - \kappa_{vs} \frac{ds}{s}$$

To eliminate singularity when $s=0$, eq. [3] is sometimes replaced by

$$[4] \quad dv^e = -\kappa_{vp} \frac{dp}{p} - \kappa_{vs} \frac{ds}{s+p_{at}}$$

where p_{at} is the atmospheric pressure.

Equations [3] and [4] appear to be stress-path independent at first glance, at least for constant κ_{vs} and κ_{vp} values. However, a closer examination reveals that both equations give inconsistent volume changes and are also stress-path de-

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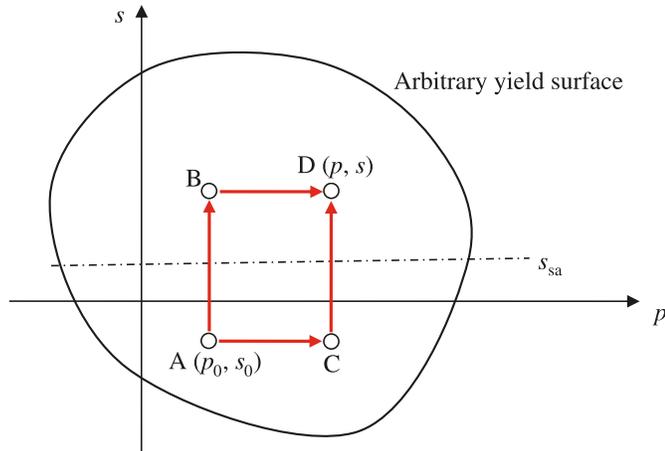
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Fig. 1. Stress paths within the elastic zone.



pendent whenever the stress state changes from saturated to unsaturated states or vice versa. Let us consider stress paths ABD and ACD in Fig. 1. For simplicity, let us assume that the pore-air pressure remains atmospheric. In this case, suction is equivalent to the negative pore-water pressure. Let us also assume that $\kappa_{vs} = \kappa_{vp}$ and the saturation suction is constant. For saturated soils, the effective stress principle holds and the elastic volume change is given by

$$[5] \quad dv^e = -\kappa_{vp} \frac{d(p+s)}{p+s} = -\kappa_{vp} \frac{dp}{p+s} - \kappa_{vp} \frac{ds}{p+s}$$

The volume change for a suction change from A to B can then be written as

$$[6] \quad \Delta v^e|_{AB} = -\kappa_{vp} \ln \frac{p_0 + s_{sa}}{p_0 + s_0} - \kappa_{vp} \ln \frac{s}{s_{sa}} \\ = -\kappa_{vp} \ln \left(\frac{p_0 + s_{sa}}{p_0 + s_0} \frac{s}{s_{sa}} \right)$$

The volume change from C to D is

$$[7] \quad \Delta v^e|_{CD} = -\kappa_{vp} \ln \frac{p + s_{sa}}{p + s_0} - \kappa_{vp} \ln \frac{s}{s_{sa}} \\ = -\kappa_{vp} \ln \left(\frac{p + s_{sa}}{p + s_0} \frac{s}{s_{sa}} \right)$$

The volume change from B to D is

$$[8] \quad \Delta v^e|_{BD} = -\kappa_{vp} \ln \frac{p}{p_0}$$

The volume change from A to C is

$$[9] \quad \Delta v^e|_{AC} = -\kappa_{vp} \ln \frac{p + s_0}{p_0 + s_0}$$

It can be seen that the elastic volume changes along ABD and ACD can be quite different. The elastic volume changes are stress-path independent only when points A, B, C, and D are all in the saturated or unsaturated zone. Equation [4] will lead to even stronger stress-path dependency due to the parameter p_{at} . It also makes the volume change due to a suction change insignificant when s is smaller than p_{at} .

Let us consider a numerical example where

$$s_{sa} = 10 \text{ kPa}, \quad p_0 = 11 \text{ kPa}, \quad s_0 = -10 \text{ kPa}, \quad p = 20 \text{ kPa}, \quad s = 20 \text{ kPa}$$

The total volume change when following stress path ABD is

$$\Delta v^e|_{ABD} = -\kappa_{vp} \ln \left(\frac{p_0 + s_{sa}}{p_0 + s_0} \frac{s}{s_{sa}} \right) - \kappa_{vp} \ln \frac{p}{p_0} = -\kappa_{vp} \ln \left(\frac{21}{1} \frac{20}{10} \frac{20}{11} \right) = -\kappa_{vp} \ln(76.4)$$

The total volume change over stress path ACD is

$$\Delta v^e|_{ACD} = -\kappa_{vp} \ln \left(\frac{p + s_{sa}}{p + s_0} \frac{s}{s_{sa}} \right) - \kappa_{vp} \ln \frac{p + s_0}{p_0 + s_0} = -\kappa_{vp} \ln \left(\frac{30}{10} \frac{20}{10} \frac{10}{1} \right) = -\kappa_{vp} \ln(60)$$

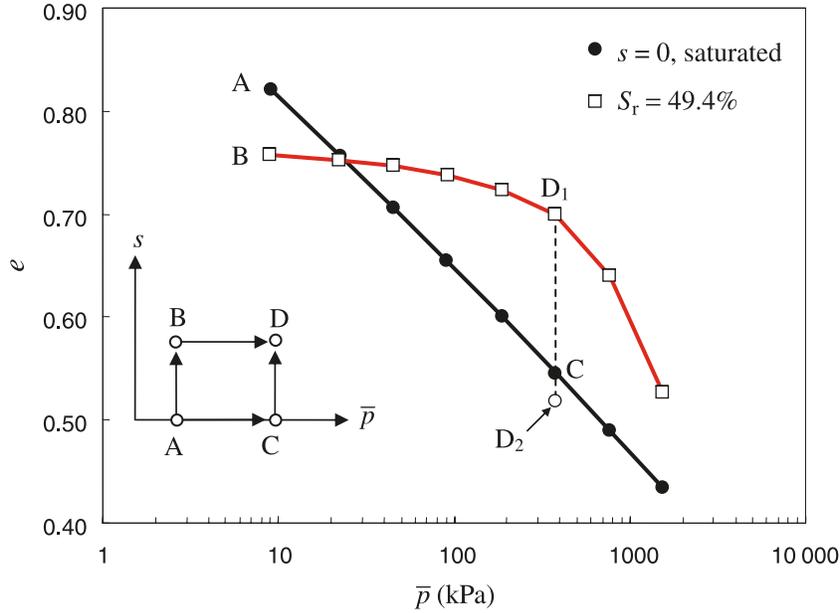
The volume change difference between the two stress paths is then $\kappa_{vp} \ln(76.4/60) = \kappa_{vp} \ln 1.27$ or about 6% of the predicted total volume change. On the other hand, the SFG model would predict a volume change difference of $\kappa_{vp} \ln(38.16/36.85) = \kappa_{vp} \ln 1.03$ or about 0.8% of the total volume change. To obtain this figure, we have to integrate the following SFG equation along the stress paths:

$$dv = -\kappa_{vp} \frac{dp}{p+s} - \kappa_{vs} \frac{ds}{p+s}$$

A semi-logarithmic SFG relationship is used to make the parameter κ_{vp} comparable with that used in eq. [3]. The volume changes along stress paths ABD and ACD are, respectively,

$$\Delta v^e|_{ABD} = -\kappa_{vp} \ln \left(\frac{p_0 + s_{sa}}{p_0 + s_0} \right) - \kappa_{vp} \frac{s_{sa} + 1}{p_0 - 1} \ln \left(\frac{s + 1}{s_{sa} + 1} \frac{p_0 + s_{sa}}{p_0 + s} \right) - \kappa_{vp} \ln \frac{p + s}{p_0 + s} = -\kappa_{vp} \ln(36.85)$$

Fig. 2. Oedometer tests on air-dry silt soil (data from Jennings and Burland 1962).



$$\Delta v^e|_{ACD} = -\kappa_{vp} \ln\left(\frac{p+s_0}{p_0+s_0}\right) - \kappa_{vp} \ln\frac{p+s_{sa}}{p+s_0} - \kappa_{vp} \frac{s_{sa}+1}{p-1} \ln\left(\frac{s+1}{s_{sa}+1} \frac{p+s_{sa}}{p+s}\right) = -\kappa_{vp} \ln(38.16)$$

Neither of the predicted volume change differences is very significant, considering the relatively small value of κ_{vp} (typically in the order of 0.01). Equations [3] or [4] are clearly not a better option than that proposed in the SFG model in terms of stress-path dependency.

A more serious problem that arises with the use of eqs. [3] and [4] is that the volume change caused by a stress change becomes undefined when points A and B (or C and D) have the same suction, but point A is inside the saturated zone and point B is inside the unsaturated zone. This inconsistency comes from the fact that neither eq. [3] nor eq. [4] recovers eq. [5] when an unsaturated soil becomes saturated. As a consequence, the stiffness matrix in the incremental stress-strain relationship becomes undefined at the transition suction between saturated and unsaturated states.

Option 2: constant κ_{vs}

Let us now consider the use of a constant κ_{vs} and let it be set to κ_{vp} in the SFG model. This alternative would lead to stress-path independency within the elastic zone. The only difference to the model would then be the integration of eq. [11] in Sheng et al. (2008). The equation would still be integrable and would result in a somewhat different yield stress (p_y). However, this alternative would not agree with the fundamental fact that changing suction under constant stress does not cause much volume change when suction is very high. This is the reason that parameter λ_{vs} must approach zero as the suction approaches 1 000 000 kPa (or even residual suction). Obviously, we cannot have a zero λ_{vs} , but a finite κ_{vs} . Such an alternative would cause a more

significant problem than that of slight stress-path dependency.

Option 3: integration according to Green’s theorem

We could integrate the following equation as suggested by the discussers to find a stress-path independent elastic model:

$$[10] \quad \frac{d\left(\frac{\kappa_{vp}}{p+s}\right)}{ds} - \frac{d\left(\frac{\kappa_{vs}}{p+s}\right)}{dp} = 0$$

The resulting κ_{vs} from such an integration would likely be a complex function of stress and suction. This function would likely be inconsistent with the basic feature of κ_{vs} (i.e., it must vary between κ_{vp} and zero as suction increases).

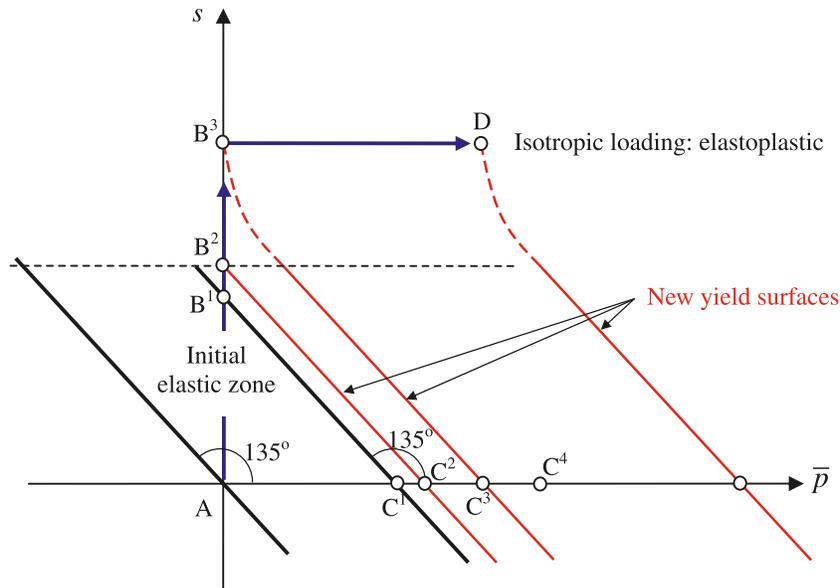
Option 4: models in terms of effective stress

Alternatively, we could use the so-called effective stress approach:

$$[11] \quad dv^e = -\kappa_{vp}(s) \frac{dp'}{p'}$$

This approach would obviously lead to a stress-path independent response in the effective stress space. However, before we start with this approach, we should first define what is meant by effective stress. Currently in the literature, all effective stress definitions for unsaturated soils involve material state variables (e.g., the degree of saturation) and can even depend on the stress path (e.g., the transition suction between saturated and unsaturated states). Therefore, the ef-

Fig. 3. Expansion of the elastic zone during drying and compression of a slurry soil.



ffective stress space where a constitutive relation is established is constantly changing with the material state or stress path. The resulting question is as follows: What is the meaning of such a constitutive relationship?

Because the so-called effective stress definition for unsaturated soils usually depends on material state or even on stress path, it is impossible to control the stress or stress path in laboratory tests. It also presents difficulties in interpreting experimental results that are obtained under prescribed net stress and suction paths.

There are also constraints on the effective stress definition that have not been thoroughly discussed in the literature. For example, the effective mean stress must decrease slower than the yield stress as suction decreases and total mean stress is kept constant to model wetting-induced collapse. On the other hand, the effective mean stress must increase faster than the yield stress as suction increases under constant total mean stress to simulate drying-induced yielding of a slurry soil. Such constraints are not always consistent with one another.

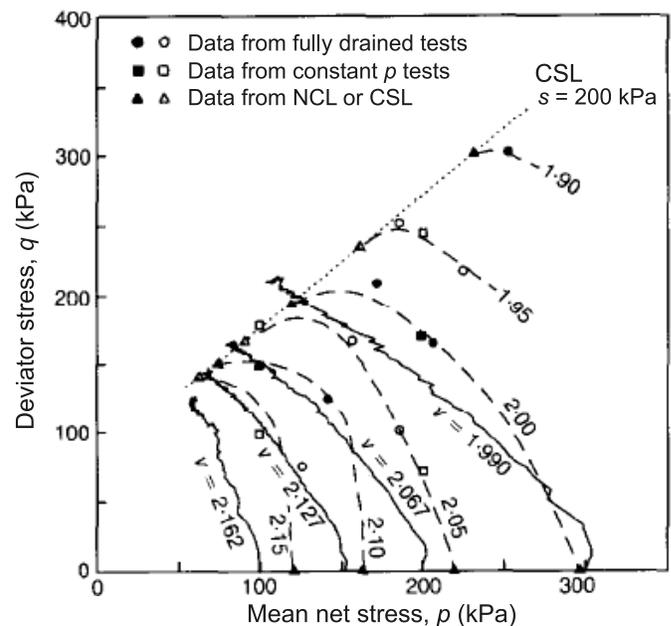
However, we can always transfer a known constitutive law established in the independent stress space into an effective stress space. Such a transfer is particularly meaningful if it simplifies the model, removes its singularities, and facilitates its implementation. For example, Sheng et al. (2003a, 2003b) have transferred the Barcelona basic model (Alonso et al. 1990) into the Bishop effective stress space:

$$[12] \quad \sigma'_{ij} = \sigma_{ij} - u_a \delta_{ij} + \chi(u_a - u_w) \delta_{ij}$$

where u_a is the pore-air pressure, u_w is the pore-water pressure, χ is the Bishop effective stress parameter, δ_{ij} is the Kronecker delta, σ_{ij} is the total stress tensor, and σ'_{ij} is the effective stress tensor.

Certain assumptions usually must be made during the transformation. For example, Sheng et al. (2003a, 2003b) used both S_r (degree of saturation) and $\sqrt{S_r}$ for the parameter χ to overcome the previously mentioned constraints. Further research is needed to transfer the SFG model into a so-called effective stress space.

Fig. 4. Specific volume contours (v) at suction of 200 kPa (Wheeler and Sivakumar 1995; reproduced with permission of Géotechnique, Vol 45, 1995). CSL, critical state line; NCL, normal compression line.

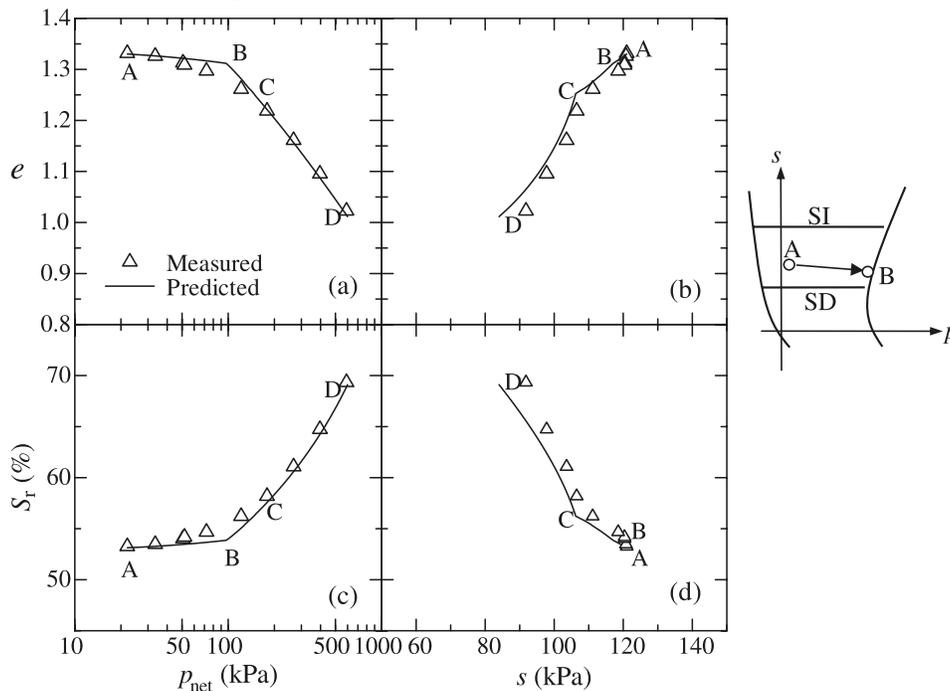


In summarizing this question, the hysteretic behaviour within the elastic zone seems to be a feature of the models that use the independent stress variables. The SFG model seems to provide one of the better solutions to the problem raised by the discussers. There may be some space for improvement. It was also stated in our paper that the equation we used is not necessarily the one that must be used.

Question 2: stress-path dependency of the elastoplastic part of the SFG model

The classical critical state models (i.e., the original and

Fig. 5. Isotropic compression test on compacted Pearl clay under undrained conditions (Sun et al. 2008; reproduced with permission of Computers and Geotechnics, Vol 35, 2008). p_{net} , net mean stress.



the modified Cam-clay models) are stress-path independent in a sense that the plastic volumetric strain for a normally consolidated soil depends only on the starting and ending stress states. However, there are numerous data supporting stress-path dependent elastoplastic behaviour (e.g., Lade and Duncan 1976; Nakai and Mihara 1984). A number of elastoplastic models have therefore been developed to tackle stress-path dependent behaviour (e.g., Nakai 1989; Nakai and Hinokio 2004; Kikumoto et al. 2007). In addition, most hypoplastic models (e.g., Kolymbas 1991) and bounding surface models (e.g., Dafalias 1986; Asaoka 2003) can deal with stress-path dependency.

For unsaturated soils, it is relatively easy to demonstrate the stress-path dependency. Jennings and Burland (1962) presented a set of oedometer test results on air-dry slurry soil, and one of their plots is shown in Fig. 2. Let us consider two stress paths ABD and ACD as shown in Fig. 2. The soil at point A is on the normal compression line for saturated states, and the stress state must be on the yield surface. Therefore, drying the soil from point A to point B and then loading it to point D will result in an elastoplastic stress path. The void ratio will change from e_A to e_B and then to e_{D1} , with e_B and e_{D1} on the normal compression line for $S_r = 49.4\%$. Alternatively, compressing the soil from point A to point C and then drying it from point C to point D will also result in an elastoplastic stress path.

The corresponding void ratio changes will follow the path ACD₂, with e_A and e_C on the normal compression line for $s = 0$ and e_{D2} below e_C . Point D₂ is estimated because drying the soil from point C to point D will generally cause some volume decrease. It is then clear that the two stress paths lead to quite different volumetric responses. More recently, Cunningham et al. (2003) presented similar results for air-dried unsaturated soils.

The crucial point here is to realize that points A, B, C,

and D are all on the current yield surfaces. A slurry soil that is air-dried and loaded according to path ABD responds in an elastoplastic manner, not just in an elastic manner. It is relatively easy to demonstrate the yield stress variation with suction for an air-dry unsaturated soil. Let us assume that the slurry soil was isotropically consolidated to point C in Fig. 3 and it has an air entry indicated by the suction at point B². Because of the effective stress principle for saturated soils, the initial elastic zone is then bounded by the two thick lines that go through points A and C¹ and are inclined to the horizontal by 45°. The yield (preconsolidation) stress at a suction corresponding to point B¹ is then zero. Drying this slurry soil under zero mean stress to a suction represented by point B² will then cause plastic yielding. The yield stress at zero suction will increase to the mean stress at point C², but the preconsolidation stress at the current suction will remain zero (point B²). Further drying will cause desaturation of the soil, but the yield stress at zero suction will generally not increase at the same rate as the suction. Let us dry the soil under zero stress to the suction at point B³. The yield stress at zero suction will then be somewhere between point C² and point C⁴ (point C³ in Fig. 3). Let us now isotropically compress the soil under the constant suction at point B³ (i.e., stress path B³D in Fig. 3). According to the data by Jennings and Burland (1962) and more recently by Cunningham et al. (2003), the isotropic compression line in the space of void ratio against logarithmic mean stress will be curved, in a pattern similar to that shown in Fig. 2 for $s > 0$. However, the isotropic compression path (B³D) is clearly elastoplastic, not purely elastic. The preconsolidation (yield) stress correspondingly increases from zero to the current stress level. Similarly, a loading path (C¹C³) followed by a drying path C³D is also elastoplastic.

We also have difficulty accepting the general statement

that the uniqueness of the state boundary surface is well accepted in the literature. The two references mentioned in the discussion both used data for compacted soils. The state boundary surface shown in Fig. 14 in Delage and Graham (1996) is for the specific stress paths: isotropic compression under constant suction followed by suction reduction under constant stress. For such a stress path, the SFG model predicts (see Fig. 3a in Sheng et al. 2008) volume changes even though the term “state boundary surface” was not used in the SFG model. The so-called experimental evidence for the uniqueness of the state boundary surface refers to Fig. 17 in Wheeler and Sivakumar (1995) and is included here as Fig. 4. It seems to us that this figure can best verify the uniqueness of the critical state line. The specific volume contours in Fig. 4 were obtained from triaxial shear tests under different stress paths but at a fixed suction value (200 kPa). It seems to us that the relationship among suction, net mean stress, and specific volume is not necessarily unique.

In summary, we do not think that the state boundary surface for a soil is necessarily unique. The stress-path dependency seems to be a good feature to have in a constitutive model for unsaturated soils, particularly when modelling unsaturated soils that are air-dried from slurry states.

Question 3: modelling hysteresis of the SWCC

This question is relevant and insightful. We agree that the hysteretic SWCCs should depend on the void ratio of the soil and consequently on the stress state. The hydraulic model used in the SFG model follows that in Sheng et al. (2004) and represents a simplification of real soil behaviour.

When an unsaturated soil specimen is sheared under undrained conditions (i.e., with constant water content), the degree of saturation can increase slightly due to the volume contraction of the soil. This has been observed in laboratory tests (e.g., Sun et al. 2008). However, the recoverable portion of the volume change is generally small, if at all. In addition, the initial state is usually within the SI and SD surfaces, not on the SD surface. To reach the SD surface, the soil must be highly wetted. Both the main drying and main wetting curves refer to drained conditions. Therefore, the elastic path (if any) AB is likely to be inside the SI and SD surfaces. In such a case, both the volume change and the change of degree of saturation are recoverable. We note that, once the volume change becomes irrecoverable, the change of degree of saturation will also be irrecoverable.

The question raised by the discussers may also apply to the compaction process (or isotropic compression) of the soil under constant water content conditions. Again, the SFG model can only model an elastic path AB that is within the SI and SD surfaces if the volume change is recoverable. Figure 5 shows typical responses of Pearl clay during undrained isotropic compression. It is clear that the changes in suction and degree of saturation are quite small when the mean stress changes from A to B and causes only elastic volume change.

Although the discussers may have amplified some effects that we feel are secondary, we note that the issue raised by the discussers indeed constitutes a theoretical challenge to the SFG model. We also note that the case illustrated in the

discussion cannot easily be modelled, even if we incorporate density-dependent SWCCs because the SWCCs usually shift to higher suctions with a decrease in void ratio. The issue definitely warrants further research.

Conclusion

In summary, we do not think that mathematical constraints should be imposed on the parameters κ_{vs} and κ_{vp} or λ_{vs} and λ_{vp} to make the volumetric model stress-path independent. Rather, these parameters should be based on the basic features of soil behaviour.

References

- Alonso, E.E., Gens, A., and Josa, A. 1990. A constitutive model for partially saturated soils. *Geotechnique*, **40**(3): 405–430.
- Asoaka, A. 2003. Consolidation of clay and compaction of sand, an elastoplastic description. *In Proceedings of the 12th Asian Regional Conference on Soil Mechanics and Geotechnical Engineering*, Singapore, 4–8 August 2003. *Edited by* C.F. Leung, K.K. Phoon, Y.K. Chow, K.Y. Yong, and C.I. Teh. World Scientific Publishing Co., Singapore. Vol. 2, pp. 1157–1196.
- Cunningham, M.R., Ridley, A.M., Dineen, K., and Burland, J.B. 2003. The mechanical behaviour of a reconstituted unsaturated silty clay. *Geotechnique*, **53**: 183–194. doi:10.1680/geot.53.2.183.37266.
- Dafalias, Y.F. 1986. Bounding surface plasticity. I: mathematical foundation and hypoelasticity. *Journal of Engineering Mechanics*, ASCE, **112**: 966–987.
- Delage, P., and Graham, J. 1996. Mechanical behaviour of unsaturated soils: understanding the behaviour of unsaturated soils requires reliable conceptual models. *In Proceedings of the 1st International Conference on Unsaturated Soils*, Paris, 6–8 September 1995. A.A. Balkema, Rotterdam, the Netherlands. Vol. 3, pp. 1223–1256.
- Jennings, J.E.B., and Burland, J.B. 1962. Limitations to the use of effective stresses in partly saturated soils. *Geotechnique*, **12**(1): 125–144.
- Kikumoto, M., Kyokawa, H., Nakai, T., Zhang, F., and Hinokio, M. 2007. Description of induced anisotropy of soils using isotropic hardening rule in modified stress space. *In Numerical Models in Geomechanics: Proceedings of the 10th International Symposium on Numerical Models in Geomechanics*. *Edited by* G.N. Pande and S. Pietruszczak. Taylor and Francis, London, UK. pp. 85–91.
- Kolymbas, D. 1991. An outline of hypoplasticity. *Archive of Applied Mechanics*, **61**: 143–151.
- Lade, P.V., and Duncan, J.M. 1976. Stress path dependent behavior of cohesion less soil. *Journal of the Geotechnical Engineering Division*, ASCE, **102**: 51–68.
- Nakai, T. 1989. An isotropic hardening elastoplastic model for sand considering the stress path dependency in three-dimensional stresses. *Soils and Foundations*, **29**: 119–137.
- Nakai, T., and Hinokio, M. 2004. A simple elastoplastic model for normally and over consolidated soils with unified material parameters. *Soils and Foundations*, **44**: 53–70.
- Nakai, T., and Mihara, Y. 1984. A new mechanics quantity for soils and its application to elastoplastic constitutive models. *Soils and Foundations*, **24**: 82–94.
- Sheng, D., Sloan, S.W., Gens, A., and Smith, D.W. 2003a. Finite element formulation and algorithms for unsaturated soils. Part I: Theory. *International Journal for Numerical and Analytical Methods in Geomechanics*, **27**: 745–765. doi:10.1002/nag.295.
- Sheng, D., Smith, D.W., Sloan, S.W., and Gens, A. 2003b. Finite

- element formulation and algorithms for unsaturated soils. Part II: Verification and application. *International Journal for Numerical and Analytical Methods in Geomechanics*, **27**: 767–790. doi:10.1002/nag.296.
- Sheng, D., Sloan, S.W., and Gens, A. 2004. A constitutive model for unsaturated soils: thermo-mechanical and algorithmic aspects. *Computational Mechanics*, **33**: 453–465. doi:10.1007/s00466-003-0545-x.
- Sheng, D., Fredlund, D.G., and Gens, A. 2008. A new modelling approach for unsaturated soils using independent stress variables. *Canadian Geotechnical Journal*, **45**(4): 511–534. doi:10.1139/T07-112.
- Sun, D.A., Sheng, D., Xiang, L., and Sloan, S.W. 2008. Elastoplastic prediction of hydro-mechanical behaviour of unsaturated soils under undrained conditions. *Computers and Geotechnics*, **35**(6): 845–852.
- Wheeler, S.J., and Sivakumar, V. 1995. An elasto-plastic critical state framework for unsaturated soil. *Géotechnique*, **45**(1): 35–53.
- Zhang, X., and Lytton, R.L. 2007. Use of state surface approach to explain the Barcelona Basic Model. *In Theoretical and numerical unsaturated soil mechanics, Edited by T Schanz*. Springer, Weimar, Germany. pp. 101–107.